

Infrared and Raman spectroscopic characterization of the silicate mineral olmiite $\text{CaMn}^{2+}[\text{SiO}_3(\text{OH})](\text{OH})$ – implications for the molecular structure



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HIGHLIGHTS

- We have studied the structure of the mineral olmiite $\text{CaMn}[\text{SiO}_3(\text{OH})](\text{OH})$.
- It forms a series with its calcium analogue poldervaartite $\text{CaCa}[\text{SiO}_3(\text{OH})](\text{OH})$.
- We used scanning electron microscopy, thermogravimetric analysis, Raman and infrared spectroscopy.
- Thermogravimetric analysis proves the mineral decomposes at 502 °C with a mass loss of 8.8%.
- Vibrational spectroscopy enables a detailed assessment of the molecular structure of olmiite.

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ABSTRACT

We have studied the mineral olmiite $\text{CaMn}[\text{SiO}_3(\text{OH})](\text{OH})$ which forms a series with its calcium analogue poldervaartite $\text{CaCa}[\text{SiO}_3(\text{OH})](\text{OH})$ using a range of techniques including scanning electron microscopy, thermogravimetric analysis, Raman and infrared spectroscopy. Chemical analysis shows the mineral is pure and contains only calcium and manganese in the formula. Thermogravimetric analysis proves the mineral decomposes at 502 °C with a mass loss of 8.8% compared with the theoretical mass loss of 8.737%. A strong Raman band at 853 cm^{-1} is assigned to the SiO stretching vibration of the $\text{SiO}_3(\text{OH})$ units. Two Raman bands at 914 and 953 cm^{-1} are attributed to the antisymmetric vibrations. Two intense Raman bands observed at 3511 and 3550 cm^{-1} are assigned to the OH stretching vibration of the $\text{SiO}_3(\text{OH})$ units. The observation of multiple OH bands supports the concept of the non-equivalence of the OH units. Vibrational spectroscopy enables a detailed assessment of the molecular structure of olmiite.

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1. Introduction

The mineral olmiite, ideally $\text{CaMn}[\text{SiO}_3(\text{OH})](\text{OH})$, forms a series with poldervaartite, it is Ca analogue – $\text{CaCa}[\text{SiO}_3(\text{OH})](\text{OH})$ [1,2]. This mineral was first found in South Africa at the N'Chwaning II mine of the Kalahari manganese fields and occurs as a product of hydrothermal alteration of primary sedimentary and low-grade metamorphic ores in a temperature range from 250 °C to 400 °C. Usually olmiite occurs in associated with celestine, buftonteinite, poldervaartite, sturmanite and hematite [1,2].

Olmiite crystallizes in the orthorhombic crystal system, with unit cell parameters $a = 9.243(3)$, $b = 9.076(9)$, $c = 10.342(9)$ Å. $V = 868(1)$ Å³ [1,2]. The atomic arrangement of olmiite is similar

to that of poldervaartite. Variations in bond distances and angles are related to the pronounced difference in the Mn content. The structure of both poldervaartite and olmiite can be schematically described, for convenience, as a sequence along [010] of alternating layers containing M2 + Si and M1 polyhedra, respectively. Nonetheless, the strong three-dimensional linkage among the polyhedral units results in a rather isotropic framework, as pointed out by Dai [1,2]. Olmiite consists of isolated $\text{SiO}_3(\text{OH})$ tetrahedra continuously connected along the [100] direction to two M_2O_7 polyhedra by sharing edges. The M2–Si chains are linked together by corner sharing to form waved sheets parallel to (010). Alternating sheets are linked together by M1 octahedra and hydrogen bonds, resulting in a strongly bonded polyhedral network [1,2].

The combination of scanning electron microscopy and vibrational spectroscopy has proven most useful for the study of the chemistry and molecular structure of minerals. In this work we

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have studied the mineral olmiite $\text{CaMn}[\text{SiO}_3(\text{OH})](\text{OH})$ using a combination of techniques.

2. Experimental

2.1. Samples description and preparation

The olmiite sample studied in this work was obtained from the collection of the Geology Department of the Federal University of Ouro Preto, Minas Gerais, Brazil, with sample code SAB-098. The sample is from the type locality in N' Chwaning II mine of the Kalahari manganese fields, South Africa. The mineral occurs in association with calcite and manganese oxides not determinate. The sample was gently crushed and the associated minerals were removed under a stereomicroscope Leica MZ4. Scanning electron microscopy (SEM) was applied to support the chemical characterization.

2.2. Scanning electron microscopy (SEM)

Experiments and analyses involving electron microscopy were performed in the Center of Microscopy of the Universidade Federal de Minas Gerais, Belo Horizonte, Minas Gerais, Brazil (<http://www.microscopia.ufmg.br>).

Olmiite crystal aggregate was coated with a 5 nm layer of evaporated Au. Secondary Electron and Backscattering Electron images were obtained using a JEOL JSM-6360LV equipment. Qualitative and semi-quantitative chemical analyses in the EDS mode were performed with a ThermoNORAN spectrometer model Quest and was applied to support the mineral characterization.

3. Electron microprobe analysis (EMP)

Chemical characterization of olmiite was carried via EMP. One single crystal was selected and prepared in epoxy resin. The single crystal was analyzed with the performance of ten spots. The chemical analysis was carried out with a Jeol JXA8900R spectrometer from the Physics Department of the Federal University of Minas Gerais, Belo Horizonte. For each selected element, we used the following standards: Ca – anortite, Si – quartz, Mg – MgO, Fe – siderite and Mn – rhodonite. H_2O was calculated by stoichiometry. The epoxy embedded olmiite sample was coated with a thin layer of evaporated carbon. The electron probe microanalysis in the WDS (wavelength dispersive spectrometer) mode was obtained at 15 kV accelerating voltage and beam current of 10 nA. Chemical formula was calculated on the basis of five oxygen atoms (O). ZAF correction was applied with support of the JEOL software XM-97312. The error in measurements is considered ± 2 wt%.

3.1. Thermogravimetric analysis – TG/DTG

TG/DTG analysis of the olmiite were obtained by using TA Instruments Inc. Q50 high-resolution TGA operating at a $10^\circ\text{C}/\text{min}$ ramp with data sample interval of 0.50 s/pt from room temperature to 1000°C in a high-purity flowing nitrogen atmosphere ($100\text{ cm}^3/\text{min}$). A total mass of 40.00 mg of finely ground samples were heated in an open platinum crucible.

3.2. Raman microprobe spectroscopy

Crystals of olmiite were placed on a polished metal surface on the stage of an Olympus BHSM microscope, which is equipped with $10\times$, $20\times$, and $50\times$ objectives. The microscope is part of a Renishaw 1000 Raman microscope system, which also includes a monochromator, a filter system and a CCD detector (1024 pixels).

The Raman spectra were excited by a Spectra-Physics model 127 He–Ne laser producing highly polarized light at 633 nm and collected at a nominal resolution of 2 cm^{-1} and a precision of $\pm 1\text{ cm}^{-1}$ in the range between 200 and 4000 cm^{-1} . Repeated acquisitions on the crystals using the highest magnification ($50\times$) were accumulated to improve the signal to noise ratio of the spectra. Raman Spectra were calibrated using the 520.5 cm^{-1} line of a silicon wafer. The Raman spectrum of at least 10 crystals was collected to ensure the consistency of the spectra.

3.3. Infrared spectroscopy

Infrared spectra were obtained using a Nicolet Nexus 870 FTIR spectrometer with a smart endurance single bounce diamond ATR cell. Spectra over the $4000\text{--}525\text{ cm}^{-1}$ range were obtained by the co-addition of 128 scans with a resolution of 4 cm^{-1} and a mirror velocity of 0.6329 cm/s . Spectra were co-added to improve the signal to noise ratio.

Spectral manipulation such as baseline correction/adjustment and smoothing were performed using the Spectralcalc software package GRAMS (Galactic Industries Corporation, NH, USA). Band component analysis was undertaken using the Jandel 'Peakfit' software package that enabled the type of fitting function to be selected and allows specific parameters to be fixed or varied accordingly. Band fitting was done using a Lorentzian-Gaussian cross-product function with the minimum number of component bands used for the fitting process. The Gaussian-Lorentzian ratio was maintained at values greater than 0.7 and fitting was undertaken until reproducible results were obtained with squared correlations of r^2 greater than 0.995.

4. Results and discussion

4.1. Chemical characterization

The SEM image of olmiite sample studied in this work is shown in Fig. 1. The image shows a olmiite crystal aggregate. The mineral occurs in association with small amounts of manganese oxides. Qualitative chemical analysis shows a Ca and Mn silicate (Fig. 2). No other elements were detected (see Table 1).

4.1.1. Thermogravimetric analysis

The thermogravimetric analysis of olmiite is recorded in Fig. 3. A mass loss of 8.8% is observed at 502°C . The theoretical mass loss based upon the formula $\text{CaMn}[\text{SiO}_3(\text{OH})](\text{OH})$ is 8.737%. The temperature range for this mass loss is very narrow. The following

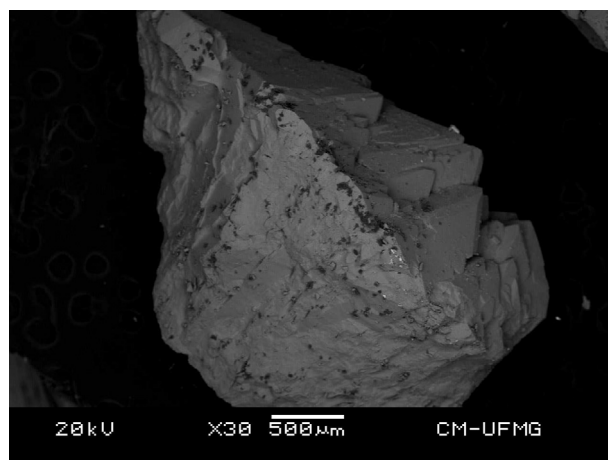


Fig. 1. Backscattered electron image (BSI) of an olmiite crystal aggregate up to 3.0 mm in length.

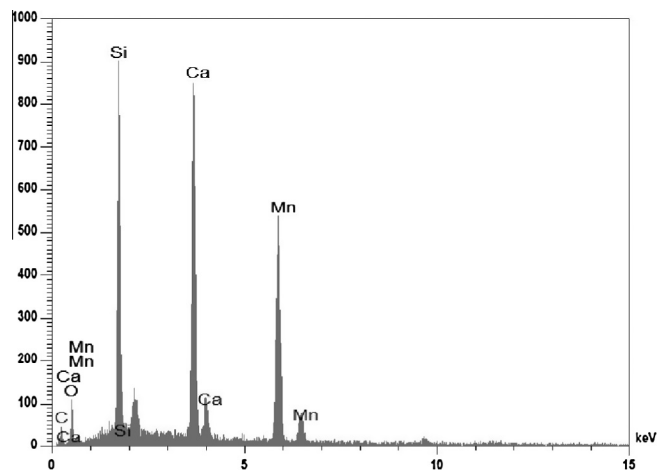


Fig. 2. EDS analysis of olmiite.

Table 1
Chemical analysis of olmiite.

| Constituent | wt.% | Number of atoms | Probe standard/crystal |
|------------------|-------|-----------------|------------------------|
| SiO ₂ | 30.39 | 1.03 | Quartz |
| CaO | 32.63 | 1.19 | Anortite |
| MnO | 27.74 | 0.77 | Rodhonite |
| MgO | 0.01 | 0.00 | MgO |
| FeO | 0.04 | 0.01 | Siderite |
| H ₂ O | 8.87 | 1.96 | Stoichiometry |
| Total | 99.02 | 4.96 | |

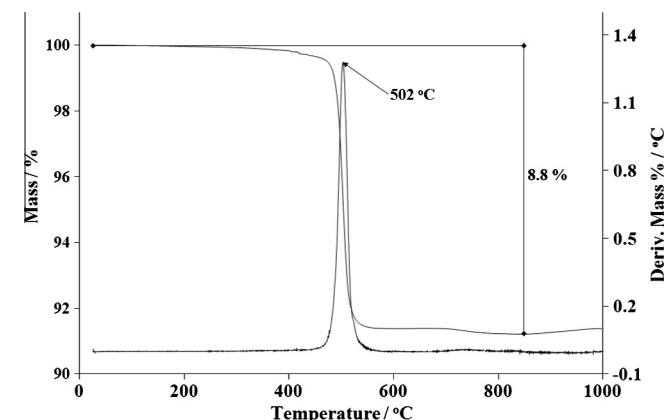
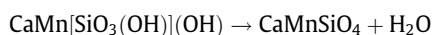


Fig. 3. Thermogravimetric and derivative thermogravimetric analysis of olmiite.

chemical reaction is proposed for the thermal decomposition of olmiite:



4.2. Vibrational spectroscopy

The Raman spectrum of olmiite over the 100–4000 cm⁻¹ spectral range is illustrated in Fig. 4a. This figure shows the position and relative intensity of the Raman bands. It is noted that there are large parts of the spectrum where no intensity is found; therefore, the Raman spectrum is subdivided into sections depending upon the type of vibration being analyzed. The infrared spectrum of olmiite over the 500–4000 cm⁻¹ spectral range is displayed in Fig. 4b. This figure shows the position and relative intensity of

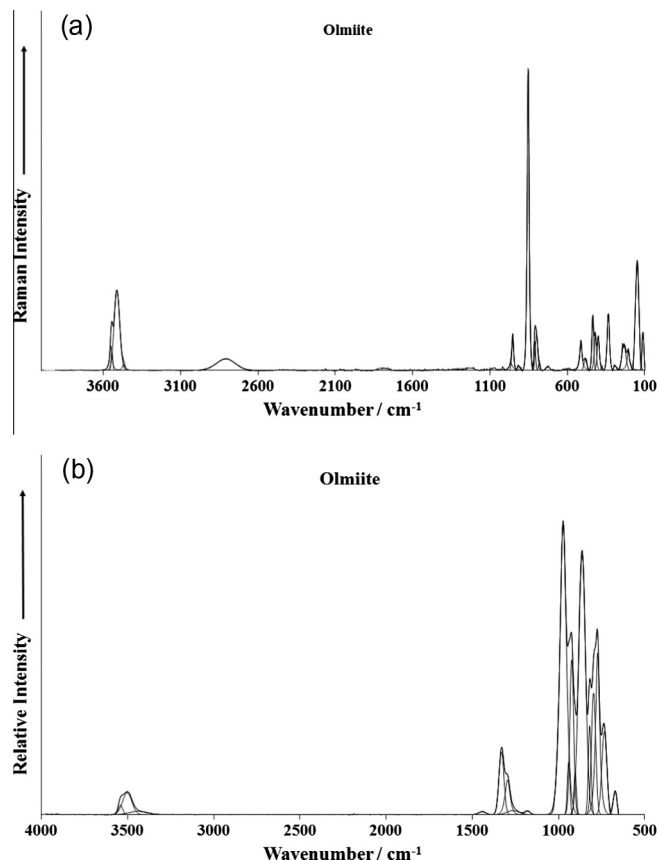


Fig. 4. (a) Raman spectrum of olmiite over the 4000–100 cm⁻¹ spectral range and (b) Infrared spectrum of olmiite over the 4000–500 cm⁻¹ spectral range.

the infrared bands. Again, there are large parts of the infrared spectrum where little or no intensity is observed and therefore, the spectrum is subdivided into sections based upon the type of vibration being studied.

The Raman spectrum of olmiite over the 700–1000 cm⁻¹ spectral range is shown in Fig. 5a. The Raman spectrum is dominated by an intense peak at 853 cm⁻¹. Dowty showed that the –SiO₃ units had a unique band position of 980 cm⁻¹ [3] (see Figs. 2 and 4 of this reference). Dowty also showed that Si₂O₅ units had a Raman peak at around 1100 cm⁻¹. Olmiite consists of isolated SiO₃(–OH) tetrahedra continuously connected along the [100] direction to two M₂O₇ polyhedra by sharing edges. The M₂–Si chains are linked together by corner sharing to form waved sheets parallel to (010). Alternating sheets are linked together by M1 octahedra and hydrogen bonds, resulting in a strongly bonded polyhedral network [1,2]. Thus, the Raman peak at 853 cm⁻¹ is assigned to the SiO stretching vibration of the SiO₃(OH) units. The two bands at 914 and 953 cm⁻¹ with a low intensity shoulder at 964 cm⁻¹ are attributed to the antisymmetric vibrations of the SiO₃(OH) units. The two Raman bands at 799 and 811 cm⁻¹ are attributed to the SiOH deformation modes.

The infrared spectrum of olmiite over the 600–1100 cm⁻¹ spectral range is shown in Fig. 5b. Strong infrared bands are observed at 860 and 970 cm⁻¹. This latter band together with the infrared bands at 900, 920 and 938 cm⁻¹ are assigned to the SiO antisymmetric stretching vibration. The infrared band at 860 cm⁻¹ is ascribed to the SiO symmetric stretching band. Other infrared bands are observed at 668, 731, 769, 793 and 816 cm⁻¹. If we follow the assignment of the Raman bands then these bands are due to hydroxyl deformation modes.

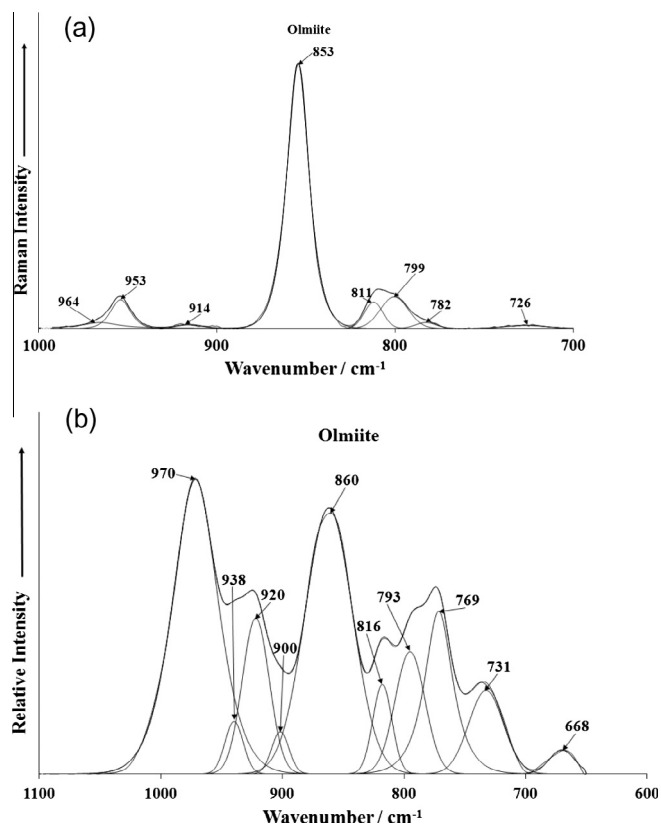


Fig. 5. (a) Raman spectrum of olmiite over the 1000–700 cm^{-1} spectral range and (b) Infrared spectrum of olmiite over the 1100–600 cm^{-1} spectral range.

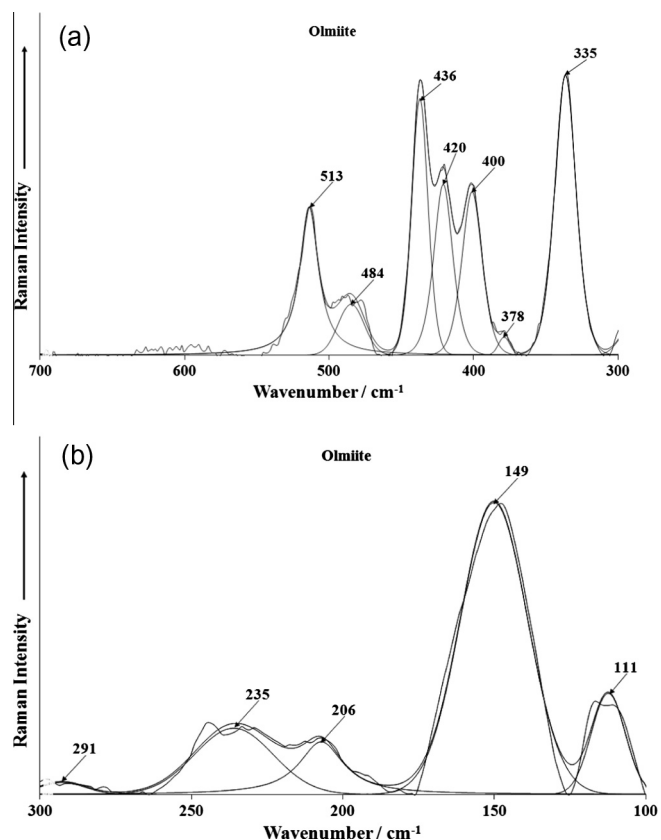


Fig. 6. (a) Raman spectrum of olmiite over the 700–300 cm^{-1} spectral range and (b) Raman spectrum of olmiite over the 300–100 cm^{-1} spectral range.

The Raman spectra over the 300–700 cm^{-1} spectral range and over the 100–300 cm^{-1} spectral range are presented in Fig. 6. Raman bands are noted at (a) 484 and 513 cm^{-1} , (b) 400, 420 and 436 cm^{-1} and (c) at 335 cm^{-1} . The first two sets of bands are assigned to OSiO bending modes. The band at 335 cm^{-1} is assigned to CaO stretching vibrations. Dowty calculated the band position of these bending modes for different siloxane units [3]. Dowty demonstrated the band position of the bending modes for SiO_3 units at around 650 cm^{-1} . This calculated value is in harmony with the higher wavenumber band observed at 663 cm^{-1} . According to Adams et al. [4] the band at 432 cm^{-1} is due to the coincidence of both the B_{2g} and E_g modes. Raman bands are observed in the

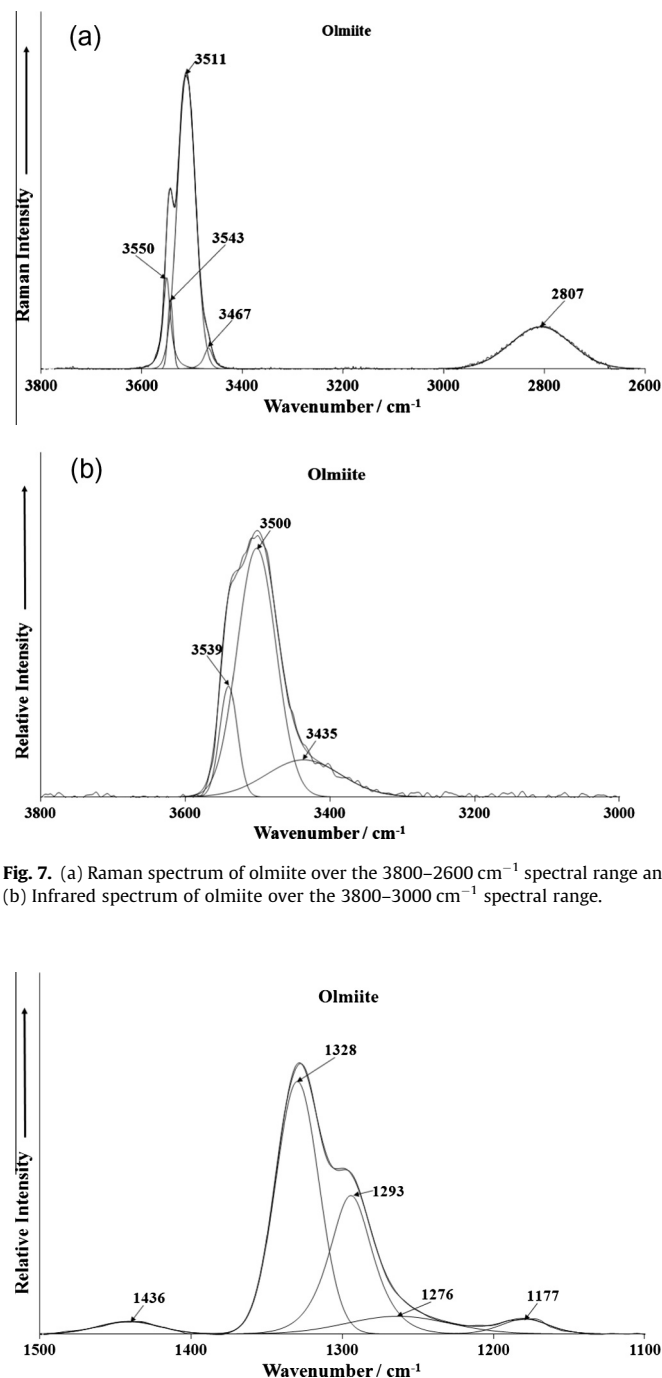


Fig. 7. (a) Raman spectrum of olmiite over the 3800–2600 cm^{-1} spectral range and (b) Infrared spectrum of olmiite over the 3800–3000 cm^{-1} spectral range.

Fig. 8. Infrared spectrum of olmiite (lower spectrum) in the 1500–1100 cm^{-1} spectral range.

far low wavenumber region at 111, 149, 296, 235 and 291 cm^{-1} . These bands are simply described as lattice vibrations.

The Raman spectrum of olmiite over the 2600–3800 cm^{-1} spectral range is reported in Fig. 7a. Two prominent peaks are observed at 3511 and 3550 cm^{-1} . These bands are assigned to the OH stretching vibration of the $\text{SiO}_3(\text{OH})$ units. The observation of multiple bands proves that the OHs are not equivalent in the olmiite structure. The infrared spectrum of olmiite over the 3000–3800 cm^{-1} spectral range is shown in Fig. 7b. Infrared bands are found at 3500 and 3539 cm^{-1} with a broad shoulder at 3435 cm^{-1} . The infrared bands support the concept of the non-equivalence of the OH units in the olmiite structure. The broad feature may be ascribed to water stretching modes; however, no water bending mode was observed in the infrared spectrum in the 1500–1800 cm^{-1} spectral range. Some infrared bands were observed in the 1100 to 1500 cm^{-1} spectral range (Fig. 8). Two infrared bands are observed at 1293 and 1328 cm^{-1} with lower intensity bands observed at 1177, 1276 and 1436 cm^{-1} . These bands are all related to SiO vibrations.

5. Conclusions

We have studied the mineral olmiite $\text{CaMn}[\text{SiO}_3(\text{OH})](\text{OH})$ which forms a continuous series with its calcium analogue poldervaartite $\text{CaCa}[\text{SiO}_3(\text{OH})](\text{OH})$ using a combination of electron microscopy with EDX analysis, thermogravimetric analysis, infrared and Raman spectroscopy. This mineral occurs as a product of hydrothermal alteration of primary sedimentary and low-grade metamorphic ores in a temperature range from 250 °C to 400 °C. These temperatures are well below the thermal decomposition of olmiite of 502 °C.

Raman spectroscopy identifies a peak at 853 cm^{-1} assigned to the SiO stretching vibration of the $\text{SiO}_3(\text{OH})$ units. Strong infrared bands observed at 860 and 970 cm^{-1} are assigned to this vibration. Two prominent Raman peaks observed at 3511 and 3550 cm^{-1} are assigned to the OH stretching vibration of the $\text{SiO}_3(\text{OH})$ units. The equivalent infrared bands found at 3500 and 3539 cm^{-1} supports the concept of the non-equivalence of the OH units in the olmiite structure. Vibrational spectroscopy enables a detailed assessment of the molecular structure of olmiite to be attained.

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